## CONTRIBUTION IN THE KNOWLEDGE OF THE NEOTECTONIC STRUCTURE OF THE GEROLIMENAS-MEZAPOU PENINSULA (MANI)

## N. Fytrolakis

Nat. Technical University, Dept. of Mining and Metallurgical Engineering, Section of Geological Sciences, 42, Patission Str., 10682 Athens, Greece

The Gerolimena-Mezapou peninsula is formed by five main and two secondary fault zones and by some smaller faults. Each fault zone is formed by normal faults, the strike of which, variates between NNW-SSE to NNE-SSW. These fault zones were formed with the begining of the uplift of the area and were reactivated in at least three main phases until the end of the uplift (lower Pleistocene-today). The last activation occured during historical times. Except the Vertical movements, the tectonic blocks are rotating about an east-west axis. So the two western tectonic blocks tilt more to the north and the two eastern to the south.

## GEOLOGICAL STRUCTURE OF SW ARGOLIS (PELOPONNESUS, GREECE)

## P. Gaitanakis, A. Photiades

IGME, Branch of Peloponnesos, 21 Ag. Vassilios Sq., 22100 Tripolis, Greece

The geological structure of SW Argolis is devided into three different geological units which are from N (Irla-Didyma areas) to S (Kranidi, Kouverta Bay areas):

a) Permo-Triassic volcanic olistostromatic formations with radiolarian cherts and middle-upper Triassic redish limestones (Vourlia Bay) overlain by upper Triassic-Jurassic limestones of Pantocrator. The upper section of Pantocrator shows pelagic formations of Ammonitico-rosso as well as redish siliceous mudstones with radiolarians topped with turbiditic and olistostromatic ophiolites (lower ophiolitic unit). The above formation is well developed East of the Ira-Didyma area.

b) Paleozoic formations (crystalline dolomites, quartzites, quartz conglomerates with characteristic grain size distribution) and Permo(?) – Triassic volcanic olistostromes (Kilada), Unconformable on the previous formations are deposited banded upper Jurassic limestones with chert intercalations which in the upper parts show laterites (pisoliths). Laterally and in places there are ophiolitic olistostromes (lower ophiolitic unit) topped, by transgression of lower Cretaceous (Aptian-Albian) with ophiolitic detritus and Upper Cretaceous greylight grey neritic limestones and Maestrichtian redish pelagic limestones overlain by Paleocenic redish marly pelites and Eocene flysch.

This formation is characteristic of the S. Fourni, Prophitis Ilias, N. Ermioni and Kouverta Bay area as a tectonic window, with especially its upper members might be correlated with