# **CONTINUOUS EXTRA-FRAMEWORK Na<sup>+</sup> RELEASE FROM** GREEK ANALCIME-RICH VOLCANICLASTIC ROCKS **ON EXCHANGE WITH NH4<sup>+</sup>**

Kantiranis N.<sup>1</sup>, Sikalidis C.<sup>2</sup>, Papastergios G.<sup>1</sup>, Squires C.<sup>1</sup> and Filippidis A.<sup>1</sup>

<sup>1</sup>Department of Mineralogy-Petrology-Economic Geology, Aristotle University, 54124 Thessaloniki, Greece, kantira@geo.auth.gr, gpapaste@geo.auth.gr,anestis@geo.auth.gr

<sup>2</sup> Department of Chemical Engineering, Aristotle University, 54124 Thessaloniki, Greece, sikalidi@auth.gr

Abstract: The continuous extra-framework cations release from Greek analcime-rich rock sample was studied, upon ammonium acetate exchange experiments (agitation time 0.25-720 hours), using atomic absorption spectrometry. The analcime-rich material was examined by X-ray powder diffraction, scanning electron microscopy equipped with energy dispersive micro-analytical system and atomic absorption spectrometry. Its sorption ability was measured using the Ammonium Acetate Saturation method. The monovalent cations  $K^+$  and  $Na^+$  after 720 hours, show only 7 and 10% of exchange, respectively. No steady state achieved for Na<sup>+</sup>. The bivalent cations Ca<sup>2+</sup> and Mg<sup>2+</sup> show better exchange, 97% for  $Ca^{2+}$  and 62% for Mg<sup>2+</sup>. The calculated rate of ion-exchange was 0.01 ppm/h for K<sup>+</sup> and Mg<sup>2+</sup>, while 0.13 ppm/h for Na<sup>+</sup> and Ca<sup>2+</sup>. The recorded behaviour on the multi-component ion-exchange system and the linear release of  $Na^+$  over  $NH_4^+$  observed at a slow rate of ion-exchange, can allow us to propose studied analcime-rich rock as a potential material for waste-water purification and pet litter.

Keywords: tuffs, cation exchange capacity, selectivity, potential uses.

## 1. Introduction

The analcime structure is made up of linked tetrahedral units proscribing three on-connected channels (Szostak, 1992). Analcime has a relatively compact structure compared with other zeolites has idealized cell and an unit of Na<sub>16</sub>(Al<sub>16</sub>Si<sub>32</sub>O<sub>96</sub>)16H<sub>2</sub>O (e.g., Meier et al., 1996; Baerlocher et al., 2001). The cubic unit cell is composed of four-, six-, and eight-membered oxygen rings that form three non-intersecting channels with pore openings of 2.6 Å. These channels encompass 24 small cavities, of which 16 are occupied by sodium cations (Saha, 1959). Analcime has maximum cation exchange capacity 4.9 meq  $g^{-1}$ (Sherman, 1978). Small pore zeolites, such as analcime, can constitute a basis for possible future technical innovations in selective adsorption and heterogeneous catalysis (Cheng et al., 2000).

High quality natural zeolites have many applications in industry, agriculture and the environment (e.g., Mumpton, 1978; Pond and Mumpton, 1984; Dyer, 1988; Tsitsishvili et al., 1992; Misaelides et al., 1993; Holmes, 1994; Ming and Mumpton, 1995; Misaelides et al., 1995a,b; Filippidis et al.,

1996; Godelitsas et al., 1996a,b; Bish and Mind, 2001). The great majority of these applications employs their cation exchange capabilities and involves the replacement of the existing extraframework cations with other cations from the surrounding environment. Due to the favourable ionexchange selectivity of natural zeolites for certain cations, these minerals have been studied for potential use in the treatment of nuclear, municipal and industrial wastewaters and acid mine drainage waters. It is of economic importance that natural zeolites are used in their natural state so that expensive purification is avoided.

The aim of this work is to study the outgoing cations (Na<sup>+</sup>, K<sup>+</sup>, Ca<sup>2+</sup> and Mg<sup>2+</sup>) of analcime-rich volcaniclastic rocks during cation exchange with NH4<sup>+</sup>. The NH4<sup>+</sup> was selected as the cation of choice as the affinity of zeolites for this cation has led to it being exploited for multiple applications such as in waste water treatment, pet litter and aquaculture. It is also used in the recognized AMmonium Acetate Saturation (AMAS) method (Bain and Smith, 1987) for the determination of the sorp-

tion ability of zeolites. Studies to evaluate how a natural zeolite exchanges with a single-cation or a multi-cation solution often use the Na- exchanged form of the zeolite, which has been shown to lead to higher cation exchange capacities (Yang et al., 1997). Results of these studies however are not applicable where the zeolite will be used in its natural form. Also, most published ion-exchange data on natural zeolites involve the exchange of two cations only. Little attention has been paid to the problem of understanding multi-component ionexchange equilibrium, despite the fact that ionexchange processes in natural systems, generally involve more than two ions. In this paper, the exchange of the extra-framework cations  $Na^+$ ,  $K^+$ ,  $Ca^{2+}$  and  $Mg^{2+}$  found in the natural zeolite with a solution of ammonium acetate (NH<sub>4</sub>AcOH), as the source of NH<sub>4</sub><sup>+</sup> cations, is studied. Greek analcimerich rock was selected due to its high zeolite content in order to study its different rates of exchange, while the results of this work are directly connected to the potential applications of the particular Greek zeolite bearing-rock.

### 2. Materials and Methods

Analcime-rich (AR) volcaniclastic rock sample was taken from the Neogene sedimentary Karlovassi basin of Samos island. The basin's volcanic rocks have been extensively zeolitized to give HEU-type, analcime, chabazite, phillipsite and mordenite of which HEU-type and analcime are the most abundant (e.g., Stamatakis, 1989; Hall and Stamatakis, 1992; Pe-Piper and Tsolis-Katagas, 1991; Kantiranis et al., 2004).

The sample was studied in its bulk form, ground (<125 µm) and homogenized. Sample for mineralogical and chemical analyses was ground further in an agate mortar. Chemical analysis of sample was carried out by AAS on a Perkin Elmer 5000 spectrometer equipped with a graphite furnace. Electron probe microanalysis of the studied zeolite phase was performed on polished thin sections of the zeolite-rich rocks by SEM-EDS (Jeol JSM-840: A scanning electron microscope, equipped with a LINK 10000 AN Energy Dispersion Analyser). Corrections were made using the ZAF-4/FLS software provided by LINK. To minimize volatilization of alkalis in the zeolites framework, the electron beam spot size was enlarged and the counting time decreased. Different minerals (micas, carbonates, feldspars) and pure metals were used as probe standards.

The experimental procedure for the cation exchange experiments entailed placing 100 mg zeolite-rich sample (<125  $\mu$ m) and 10 mL of 1N ammonium acetate (of pH 7.0) solution into screw-top glass tubes. The tubes were sealed and agitated for 0.25, 0.75, 1.25, 2, 4, 6, 8, 12, 18, 24, 36, 48, 72, 96, 120, 144, 168, 192, 216, 240 and 720 h. The mixtures were then centrifuged and the liquor decanted into a 200 mL volumetric flask. The solid was then washed with 10 mL of deionised water three times and the washings were combined with the liquor and the volume adjusted to 200 mL. The concentration of Na<sup>+</sup>, K<sup>+</sup>, Ca<sup>2+</sup> and Mg<sup>2+</sup> cations in this solution was analyzed by AAS on the Perkin Elmer 5000 spectrometer.

The mineralogical composition of the sample was determined by X-ray Powder Diffraction (XRPD) method. The XRPD analysis was performed using a Philips PW1710 diffractometer with Ni-filtered CuK<sub> $\alpha$ </sub> radiation on randomly oriented samples. The counting statistics of the XRPD study were: step size: 0.01° 20, start angle: 3°, end angle: 63° and scan speed: 0.02° 20/sec. Quantitative estimates of the abundance of the mineral phases were derived from the XRPD data, using the intensity of a certain reflections and external standard mixtures of minerals.

In order to determine the sorption ability of the studied zeolite-rich sample, the sample was treated with 1M ammonium acetate (NH<sub>4</sub>OAc) aqueous solution, according to the AMAS method. Approximately 125 mg of the <125 µm sample was added to a centrifuge tube with 10 mL of the 1N NH<sub>4</sub>OAc solution. The suspension was well shaken, agitated for 24 hours, and then centrifuged. The clear liquid was discarded and the NH<sub>4</sub>OAcsaturation procedure repeated 9 times, adding fresh 10 mL of NH<sub>4</sub>OAc solution each time. After the completion of the 10-day NH<sub>4</sub>OAc-saturation, the excess NH<sub>4</sub>OAc was washed with 10 mL of 99% isopropyl alcohol, well shaken, and centrifuged. The clear supernatant liquid was discarded and the procedure repeated five times. The sample then was dried in room temperature. Following the  $NH_4OAc$  saturation, the  $NH_4^+$  ions retained by the zeolite-rich sample, are converted using a strong base to NH<sub>3</sub> and analysed by an ammonia electrode. The NH<sub>3</sub> concentration was determined using an Orion potentiometric ammonia gas electrode combined with a Jenway 3045 pH/mV/ion analyser. Each air-dried NH<sub>4</sub><sup>+</sup>-saturated sample was placed in a 100 mL Pyrex beaker containing a

Teflon covered stirring bar. Deionised nitrogenfree water (50 mL) was added and the solution was stirred to suspend the sample. The electrode was immersed into the suspension taking care to prevent entrapment of air under the concave tip. By addition of 0.5 mL of 10M NaOH, the NH<sub>3</sub> measurements were taken at constant level achievement. The electrode calibration was performed daily using ammonium calibrating solutions of 10, 100 and 1000 ppm provided by Jenway and hourly using the 10 ppm ammonium solution.

#### 3. Results and Discussion

The quantitative mineralogical composition of the studied zeolite-rich sample is shown in Table 1. Analcime, micas (muscovite and/or illite) and clay minerals (mainly smectite) constitute the microporous minerals of the zeolite-bearing rock. The non-microporous minerals found in studied sample are quartz and feldspars. This classification is based on their ability to sorb cations on their structure. Amorphous material (volcanic glass) was also determined in the sample (10 wt %). The analcime content was determined 71 wt %., while in minor amounts micas (5 wt %), and clay minerals (2 wt %, mainly smectite) were also found. The total microporous mineral content in sample AR was 78 wt %.

Table 1. Quantitative mineralogical composition(wt %) of the studied sample.

Sample	Α	Μ	Cl	TMM	Q	F	Am
AR	71	5	2	78	4	8	10
AR · analo	ime-ri	ch A·	Analci	me M∙M	uscov	ite and/	or illite

AR: analcime-rich, A: Analcime, M: Muscovite and/or illite, Cl: Clay minerals (mainly smectite), TMM: Total microporous minerals (Zeolite+M+Cl), Q: Quartz, F: Feldspars, Am: Amorphous.

The chemical composition of the studied zeoliterich sample is shown in Table 2. AR sample was containing 56.11 wt % SiO<sub>2</sub>, while the Al<sub>2</sub>O<sub>3</sub> content was measured 18.80 wt %. Low amounts (<0.20 wt %) of TiO<sub>2</sub> and MnO were also found, while the Fe<sub>2</sub>O<sub>3T</sub> content was measured 2.98 wt %. The total percentage of oxides of the exchangeable cations Na<sup>+</sup>, K<sup>+</sup>, Ca<sup>2+</sup> and Mg<sup>2+</sup> was 12.07 wt %, with Na<sup>+</sup> being the main exchangeable cation. Loss on ignition was measured 9.43 wt %.

In Table 3 the measured sorption ability

(meq/100g), the theoretical and the expected C.E.C. (meq/100g) of the studied zeolite-rich sample are given. The measured sorption ability was derived experimentally by the AMAS method. The theoretical C.E.C. of the sample was calculated from the quantitative mineralogical composition using values of 454 meq/100g for analcime, 20 meq/100g for muscovite-illite and 100 meq/100g for smectite (Mumpton, 1977; Deer et al., 1992). The expected C.E.C. was calculated from the total extra-framework cations of the contained zeolite, measured by SEM-EDS (Table 4). For the calculations of the theoretical and expected C.E.C. values the results were modified according to the zeolite content of the rock.

The measured sorption ability of the sample was too close to the expected C.E.C., meaning that the main process involving the sorption procedure of the studied sample was the ion-exchange.

Table 3. Values (meq/100g) of measured sorption ability, theoretical and expected C.E.C. of the studied sample.

Sample	Measured sorp- tion ability	Theoretical C.E.C. <sup><i>a</i></sup>	Expected C.E.C. <sup><i>a,b</i></sup>
AR	334	325 <sup>c</sup>	329

<sup>*a*</sup> Values modified according to TMM content (zeolite+micas+clay minerals) of the sample, <sup>*b*</sup> Measured from the zeolite micronalysis (by SEM-EDS), <sup>*c*</sup> Amorphous ability to sorb ions was ignored.

# 3.1. The ion-exchange of extra-framework cations for NH<sub>4</sub><sup>+</sup>

The cation exchange behaviour of the zeolite-rich sample was investigated in order to determine which of the exchangeable ions is released first, the rate and the total amount of exchangeable cation occurs at the steady state, and the influence of the multionic system on the exchange behaviour. The SEM-EDS analyses gave accurate zeolite's chemical formula in the sample (several readings were taken for each zeolite and the average values tabulated in Table 4) providing clear data on how many and what species of potential exchangeable cations are contained in the zeolite structure and also the number of water molecules associated with the zeolite. The positions of both extraframework-cations and water molecules depend on the nature of the exchangeable cations of

Table 2. Chemical composition (wt %) of the studied sample.

1 4010 2. 0	menneur	compor	mion (wi	70) OI the	bluarec	* Sumpre						
Sample	$SiO_2$	TiO <sub>2</sub>	$Al_2O_3$	Fe <sub>2</sub> O <sub>3T</sub>	MnO	MgO	CaO	Na <sub>2</sub> O	K <sub>2</sub> O	LOI	Total	
AR	56.11	0.19	18.80	2.98	0.10	0.50	0.94	6.36	4.27	9.43	99.68	
			0.1	(								

AR: Analcime-rich, LOI: Loss of ignition (at 1050°C for 2 hours).

the zeolites (Armbruster and Gunter, 1991; Gunter et al., 1994). However, the exact structural arrangement in the channels also depends on the specific Si/Al framework distribution and the water content, which is dependent on the water vapour pressure surrounding the sample. Thus, the exchangeable cations and channel water molecules are interdependent. Changes in cation composition cause changes in the amount and structural distribution of water molecules; likewise, changes in channel water content affect the positions of the channel cations (Bish and Boak, 2001). and  $Co^{2+}$  move into analcime.

Despite the slow exchange, the sorption ability value of 334 meq/100g obtained, using the AMAS technique, is very close to the expected value of 329 meq/100g calculated from the total percentage of extra-framework cations within the analcime structure (Tables 3 and 4). The very low degree of exchange observed in our experiment is probably due to the fact that AR zeolitic material being contacted with the same ammoniacal solution for 30 days. The AMAS method significantly improves the rate of cation exchange by replacing the solu-

Table 4. Formulae and total exchangeable cations (ppm) of the studied zeolite.

Sample	Formulae		Total exchangeable cations				
Sample	Formulae	$Na^+$	$\mathbf{K}^+$	Ca <sup>2+</sup>	$Mg^{2+}$		
AR	$(Mg_{0.20}Ca_{0.28}K_{0.58}Na_{14.97})(Fe_{0.15}Al_{15.80}Si_{31.92}O_{96})\cdot 16H_2O$	913	60	30	13		

Micas and clay minerals (mainly smectite) found in AR sample also contain alkaline and alkaline earth exchangeable cations. However, due to their small weight percentages in contrast with the zeolite content of the rock, for the purpose of this discussion it is assumed that the cations released originate from the zeolite phase only.

The kinetic curves of AR sample for cation exchange of Na<sup>+</sup>, K<sup>+</sup>, Ca<sup>2+</sup> and Mg<sup>2+</sup> were presented in Figures 1(a)-1(d). Statistical analysis of the data was obtained using Microcal Origin v.6.0 software, while statistical models and correlation coefficients (R<sup>2</sup>) were shown in each figure.

The release of Na<sup>+</sup> cations from sample AR (Fig. 1a) was linear with time over the whole 30 days (720h) period studied and a steady state was not reached. However, it is known that analcime, which has a dense structure, non-intersecting channels and small individual cavities, exhibits very limited ion exchange of its Na<sup>+</sup> for other cations. Particularly, at room temperature analcime does not readily exchange, but at elevated temperatures Na can be completely exchanged by other cations such as K<sup>+</sup> and NH<sub>4</sub><sup>+</sup> (Vaughan, 1978). Balgord and Roy (1974) reported that Na-analcime can be completely exchanged by K<sup>+</sup>, Ag<sup>+</sup>, Tl<sup>+</sup>, NH<sub>4</sub><sup>+</sup> and Rb<sup>+</sup> cations at elevated temperatures, whereas only small amounts of Sr<sup>2+</sup>, Mg<sup>2+</sup>, Ni<sup>2+</sup>

tion every 24 hours for 10 days thus encouraging cation exchange by removing any competing cations from solution.

After 30 days (Fig. 1b) the concentration of K in solution was 4 ppm (Table 5) indicating the exchange of only 7 % of K<sup>+</sup> cations found in the aluminosilicate framework (Table 4). Bivalent cations  $Ca^{2+}$  and  $Mg^{2+}$  show better exchange. The Ca<sup>2+</sup> concentration achieve a steady state of 29 ppm (Table 5) after 216 hours of agitation (Fig. 1c) indicating an almost fully exchange (97%) of  $Ca^{2+}$ cations found in analcime framework (Table 4). On the other hand,  $Mg^{2+}$  did not achieve a steady state (Fig. 1d), but after 30 days (720 h) its concentration was 8 ppm (Table 5), indicating the exchange of 62% of Mg<sup>2+</sup> cation found in the analcime framework (Table 4). No steady state achieved for Na<sup>+</sup> (Fig. 1a) and only a percentage of 10% was exchanged from the structure of analcime (Tables 4 and 5).

In Table 5 the agitation time (h), the cation concentration (ppm), the degree of cation exchange (%) and the rate of exchange (ppm/h) once a steady state has been achieved, were summarized for sample AR.

Taking into consideration all the above data, the order of cation release from AR sample, is: Ca<sup>2+</sup>

Table 5. Ion-exchange statistics for analcime-rich sample.

ruble 5. fon exchange statistics for anarchile field sample.							
Extra framework ion	$Na^+$	$\mathbf{K}^+$	Ca <sup>2+</sup>	$Mg^{2+}$			
Agitation time until steady state (h)	No steady state	290	216	720			
Concentration at steady state (ppm)	After 720 h: 94	4	29	8			
Degree of exchange (%)	10	7	97	62			
Rate of exchange (ppm/h)	0.13	0.01	0.13	0.01			

Ψηφιακή Βιβλιοθήκη Θεόφβαστος - Τμήμα Γεωλογίας. Α.Π.Θ.

 $>K^+>>Mg^{2+}>>>Na^+$ , while the order of the degree of total cation release, is:  $Ca^{2+}>Mg^{2+}>>Na^+\approx K^+$ . Furthermore, the rate of ion-exchange was calculated for K<sup>+</sup> 0.01 ppm/h, for  $Ca^{2+}$  0.13 ppm/h, for  $Mg^{2+}$  0.01 ppm/h, while Na<sup>+</sup> exchanged in a constant rate of 0.13 ppm/h (Table 5). Using this rate of ion-exchange, the order for sample AR, is: Na<sup>+</sup>=Ca<sup>2+</sup>>>K<sup>+</sup>=Mg<sup>2+</sup>. Dyer and Yusof (1989) suggest a systematic-continuous replacement of Na<sup>+</sup> by NH<sub>4</sub><sup>+</sup> in analcime structure, which reduces the water content of analcime.

The linear release of Na<sup>+</sup> over NH<sub>4</sub><sup>+</sup> observed at a slow rate of ion-exchange in our experiments, can allow us to propose studied analcime-rock as a potential material for waste-water purification and pet litter. Analcime can act as an ion sieve (Barrer, 1950; Barrer and Hinds, 1953; Ames, 1966; Balgord and Roy, 1971; Moroz et al., 1998; Seryotkin et al., 2000; Likhacheva et al., 2002), while Dyer and Yusof (1987; 1989) carried out extensive studies on cation and water self-diffusion, concluding that analcime has the potential for use in the storage and disposal of titrated water.

#### 4. Conclusions

The exchange of the extra-framework cations  $Na^+$ ,  $K^+$ ,  $Ca^{2+}$  and  $Mg^{2+}$  found in Greek zeolite-rich rock was studied in its natural form in order to determine its ion exchange behaviour. It is of great importance to understand the behaviour of a multicomponent ion-exchange system, because processes of ion removal in natural systems, generally involve more than two ions. Zeolite-rich sample has analcime content 71 wt %.  $Na^+$  was the main exchangeable cation. Loss of ignition was measured 9.43 wt %.

The exchange of initial cations occurs very slow. Monovalent cations  $K^+$  and  $Na^+$  show over the whole period of the experiments (720 hours), only 7 and 10% of exchange, respectively. No steady state achieved for  $Na^+$ . Bivalent cations  $Ca^{2+}$  and



Figure 1(a-d). Kinetic curves of (a)  $Na^+$ , (b)  $K^+$ , (c)  $Ca^{2+}$  and (d)  $Mg^{2+}$  for the AR sample.

Ψηφιακή Βιβλιοθήκη Θεόφραστας - Τμήμα Γεωλογίας. Α.Π.Θ.

 $Mg^{2+}$  show better exchange. The Ca<sup>2+</sup> was almost fully exchanged (97%), while  $Mg^{2+}$ , despite that did not achieve a steady state after the end of experiments (720 hours), shows an exchange of 62%. The rate of ion-exchange was calculated for K<sup>+</sup> 0.01 ppm/h, for Ca<sup>2+</sup> 0.13 ppm/h, and for Mg<sup>2+</sup> 0.01 ppm/h. Na<sup>+</sup> was exchanged in a continuous constant rate of 0.13 ppm/h. Thus, using the rate of ion-exchange, the order for sample AR, is: Na<sup>+</sup>≈Ca<sup>2+</sup>>>K<sup>+</sup>≈Mg<sup>2+</sup>.

The continuous release of  $Na^+$  over  $NH_4^+$  observed at a slow rate of ion-exchange in our experiments, can allow us to propose studied analcime-rich rock as a potential material for waste-water purification and pet litter.

#### References

- Ames L.L., 1966. Cation exchange properties of wairakite and analcime. Am. Mineral., 51, 903-908.
- Armbruster T., Gunter M.E., 1991. Stepwise dehydration of heulandite-clinoptilolite from Succor Creek, Oregon, U.S.A.: A single-crystal X-ray study at 100 K. Am. Mineral. 76, 1872-1883.
- Baerlocher Ch., Meier W.M., Olson D.H., 2001. Atlas of zeolite framework types, Elsevier, Amsterdam, 308 p.
- Bain C., Smith L., 1987. Chemical analysis. In: Wilson, M. (Ed.), A handbook of determinative methods in clay mineralogy. Blackie, Glasgow, pp. 248-274.
- Balgord W., Roy R., 1974. Crystal Chemical Relationships in the Analcite Family, I. Synthesis and Cation Exchange Behavior. In Molecular Sieve Zeolites-1, Adv. Chem. Ser., 101, Amer. Chem. Soc., Washington, D.C., 140-148.
- Barrer R.M., 1950. Ion-exchange and ion-sieve processes in crystalline zeolites. J. Chem. Soc. Resumed, 2342-2350.
- Barrer R.M. and Hinds L., 1953. Ion-exchange in crystals of Analcite and Leucite. J. Chem. Soc., 1879-1888.
- Bish D.L., Boak J.M., 2001. Clinoptilolite-Heulandite nomenclature. In: Bish, D.L., Ming, D.W. (Eds.), Natural Zeolites: Occurrences, Properties, Applications, Reviews in Mineralogy & Geochemistry, vol. 45. The Mineralogical Society of America, Washington DC, pp. 207-216.
- Bish D.L., Ming D.W., 2001. Natural Zeolites: Occurrences, Properties, Applications. Reviews in Mineralogy & Geochemistry, vol. 45, The Mineralogical Society of America, Washington DC, 654 p.
- Cheng X., Zhao P., Stebbins J.F., 2000. Solid state NMR study of oxygen site exchange and Al-O-Al site concentration in analcime. American Mineralogist, 85, 1030–1037.
- Deer W.A., Howie R.A., Zussman J. 1992. An Introduction to the Rock-Forming Minerals, 2<sup>nd</sup> edn. Longman, London, 712 p.
- Dyer A., 1988. An Introduction to Zeolite Molecular Sieves. J. Wiley, Chichester, 149 p.

- Dyer A., Yusof A.M., 1987. Diffusion in heteroionic analcimes: Part 1. Sodium-potassium-water system. Zeolites 7, 191-196.
- Dyer A., Yusof A.M., 1989. Diffusion in heteroionic analcimes: Part II. Diffusion of water in sodium/thallium, sodium/lithium, and sodium/ammonium analcimes. Zeolites 9, 129-135.
- Filippidis A., Godelitsas A., Charistos D., Misaelides P., Kassoli-Fournaraki A., 1996. The chemical behaviour of natural zeolites in aqueous environements: Interactions between low-silica zeolites and 1M NaCl solutions of different initial pH-values. Appl. Clay Sci. 11, 199-209.
- Godelitsas A., Misaelides P., Charistos D., Filippidis A., Anousis I., 1996a. Interaction of HEU-type zeolite crystals with Thorium aqueous solutions. Chem. Erde 56, 143-156.
- Godelitsas A., Misaelides P., Filippidis A., Charistos D., Anousis I., 1996b. Uranium sorption from aqueous solutions on sodium-form of HEU-type zeolite crystals. J. Radioan. Nucl. Ch. Ar. 208, 393-402.
- Gunter M.E., Armbruster T., Kohler T., Knowles C.R., 1994. Crystal structure and optical properties of Naand Pb-exchanged heulandite-group zeolites. Am. Mineral. 79, 675-682.
- Hall A., Stamatakis M., 1992. Ammonium in zeolitized tuffs of the Karlovassi basin, Samos, Greece. Can. Mineral. 30, 423-430.
- Holmes D., 1994. Zeolites. In: Carr, D. (senior Ed.), Industrial Minerals and Rocks, Braun-Brumfield, Inc., Ann Ardor, Michigan, pp. 1129-1158.
- Kantiranis N., Stamatakis M., Filippidis A., Squires C., 2004. The uptake ability of the clinoptilolitic tuffs of Samos island, Greece. Bull. Geol. Soc. Greece 36/1, 89-96.
- Likhacheva A.Yu., Paukshtis E.A., Seryotkin Yu.V., Shulgenko S.G., 2002. IR spectroscopic characterization of NH<sub>4</sub>-analcime. Phys. Chem. Miner., 29, 617– 623.
- Meier W.M., Olson D.H., Baerlocher Ch., 1996. Atlas of Zeolite Structure Types, fourth ed. Butterworth-Heinemann, Stoneham, MA, pp. 42–43.
- Ming D.W., Mumpton F.A., 1995. Natural zeolites '93: Occurrence, Properties, Uses. International Committee on Natural zeolites, Brockport, New York.
- Misaelides P., Godelitsas A., Filippidis A., 1995a. The use of zeolithoforous rocks from Metaxades-Thrace, Greece, for the removal of caesium from aqueous solutions. Fresen. Environ. Bull. 4, 227-231.
- Misaelides P., Godelitsas A., Filippidis A., Charistos D., Anousis I., 1995b. Thorium and uranium uptake by natural zeolitic materials. Sci. Total Environ. 173/174, 237-246.
- Misaelides P., Godelitsas A., Haristos D., Noli F., Filippidis A., Sikalidis C., 1993. Determination of heavy metal uptake by the sodium form of heulandite using radiochemical techniques. Geol. Carpath.-Ser. Clays 44(2), 115-119.
- Moroz N.K., Seryotkin Yu.V., Afanassiev I.S., Belitsky

I.A., 1998. Arrangement of the extraframework cations in  $NH_4$ -analcime. J. Struct. Chem. Engl. Tr., 39(2), 281–283.

- Mumpton F.A., 1977. Natural Zeolites. In: Mumpton, F.A. (Ed.), Mineralogy and Geology of Natural Zeolites. Mineralogical Society of America Short Course Notes, v. 4, Washington D.C., pp. 1-17.
- Mumpton F.A., 1978. Natural zeolites: A new industrial mineral commodity. In: Sand, B.L., Mumpton, F.A. (Eds.), Natural zeolites: Occurrences, Properties, Uses. Pergamon Press, New York, pp. 3-27.
- Pe-Piper G., Tsolis-Katagas P., 1991. K-rich mordenite from Late Miocene rhyolitic tuffs, Island of Samos, Greece. Clay. Clay Miner. 39, 239-247.
- Pond G.W., Mumpton F.A., 1984. Zeo-Agriculture: Use of Natural Zeolites in Agriculture and Aquaculture. Westview Press, Colorado, 296.
- Saha P., 1959. Geochemical and X-ray investigation of natural and synthetic analcites. American Mineralogist, 44, 300-313.
- Seryotkin Yu.V., Bakakin V.V., Belitsky I.A., Fursenko B.A., 2000. Ag-exchanged analcime: crystal structure and crystal chemistry. Micropor. Mesopor. Materials, 39, 265-273

- Sherman J.D., 1978. Ion exchange separations with molecular sieve zeolites. AIChE Symposium Series 74(179), 98-116.
- Szostak R., 1992. Handbook of Molecular Sieves. Van Nostrand Reinhold, New York, 569 p.
- Stamatakis M., 1989. Authigenic silicates and silica polymorphs in the Miocene saline-alkaline deposits of the Karlovassi basin, Samos Island, Greece, Econ. Geol. 84, 788-798.
- Tsitsishvili G.V., Andronikashvili T.G., Kirov G.N., Filizova L.D., 1992. Natural zeolite. Ellis Horwood, New York, 274 p.
- Vaughan D.E.W., 1978. Properties of natural zeolites. In: Sand L.B., Mumpton F.A. (Eds.), Natural zeolites, Occurrence, Properties, Uses. Permangon Press, London, pp. 353-371.
- Yang P., Stolz J., Armbruster T., Gunter M.E., 1997. Na, K, Rb, and Cs exchange in heulandite single crystals: Diffusion kinetics. Am. Mineral. 82, 517-525.

Ψηφιακή Βιβλιοθήκη Θεόφραστος - Τμήμα Γεωλογίας. Α.Π.Θ.