deposits (red shales). They pass into thin- and medium-bedded turbidites with intercalations of allodapic limestones and marls and thick-bedded turbidites. Finally, thin-bedded turbidites are deposited. The cycles reflect the main evolutionary stages of the Magura basin.

MINERALOGY AND POLLEN AND SPORE ASSEMBLAGES AT THE K/T BOUNDARY ON SEYMOUR ISLAND, ANTARCTICA

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The Cretaceous/Tertiary boundary (K/T) on Seymour Island is well known through various paleontological and sedimentological studies (see Feldmann, 1988). This constribution presentes new results on the now accepted boundary zone. They are based on mineralogical and quantitative continental palynology studies.

No obvious K/T boundary can be traced from the pollen and spores diagram: Percentages show a gradual change in the pteridophytic flora form the Maastricthian to the Paleocene. This change is attributed to a chemical reduction caused by the hydrological input form a Magellanic moorland vegetation source. Slow changes in the mineralogical composition support this edaphic iterpretation. No iridium-enriched level could be found.

Literature

FELDMANN, R.M. (1988): Geology and Paleontology of Seymour Island, Antarctic Peninsula. Geol. Soc. Am., Mem. 169, p. 567.

THE DEVELOPMENT OF THE ALBANIAN ALPS ZONE; A COMPARISON WITH THE HIGH KARST OF THE DINARIDES AND THE PARNASSUS ZONE IN GREECE

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The Albanian Alps zone is the south to southwest prolongation of the High Karst zone of the Dinarids. Its sediments are equivalent in lithology and age to those of the Parnass zone of Greece. At the end of the Paleozoic and during the Mesozoic, the Albanian Alps zone was generally a carbonate platform.

Lower Triassic sediments are transgressive and consist of terrigenous deposits such

as thick-bedded heterogeneous debris flow conglomerates (sometimes with fusulinids), sandstone, turbidites and shales. Similar to the oceanic basin of the Mirdita zone, the Albanian Alps zone was influenced by the rifting of the continental crust during Anisian - Labinian times. A pelagic facies of red nodular limestones with ammonites (Han Bulog facies) and a neritic facies of massive algal limestones and dolomites developed at these times. A bauxite horizon was formed above the limestones and dolomites. In the Upper Triassic, massive neritic limestones and dolomites containing megalodonts were deposited.

In the Liassic, a very shallow marine Lithiotis facies with abundant carbonatic nodules developed. Dogger sediments show an oolithic biosparitic microfacies which was deposited in a agitated environment. Coral-algal bioclastic limestones of the Kimmeridgian were deposited in a very shallow environment. They are followed by a red horizon of bauxitic clay. During Tithonian, laggonal algal limestones (*Clypeina jurassica* microfacies), alternating frequently with brackish to freshwater algal limestones (charophyte microfacies) were deposited.

In the Albanian Alps zone, the Tithonian - Neocomian boundary lies within shallow water sediments (*Campbelliella striata* microfacies). It is followed by sediments of *Salpingoporella annulata* microfacies. The Barremian-Aptian is characterized by bedded algal bioclastic limestones (*Salpingoporella dinarica* microfacies) alternating with nerineid and requienid beds (Urgonian facies).

Cenomanian sediments consist of foraminiferal (Alveolinidae) biomicrites. They are followed by a chondrodonta facies of Turonian age and by a Senonian rudistid facies. In the late Maastrichtian a pelagic environment (platform - internal depression of Selca) developed in the Malesis Madhe subzone of the Albanian Alps zone. Gray and red marly pelagic sediments were deposited during the Paleocene. They are followed by terrigenous flysch sediments of Eocene age.

As a result of Jurassic rifting, narrow fracture zones of greater depth delimited the platform of the Albanian Alps towards the NE, S and SW. This is the Velbone zone. It is equivalent to the Praecarstic zone of the Dinarids and is transitional between the Northern Carbonate Platform of the Albanian Alps and the Gukali basin. Its sediments are condensed pelagic limestones of Liassic and Dogger age and pelagic calpionellid limestones of the Upper Jurassic and the Neocomian. The Maastrichtian Vermushi flysch unconformably overlies deposits of different ages.