

MICROSTRUCTURAL FINITE STRAIN ANALYSIS OF THE HAFAFIT GRANITOID DOMES SOUTH CENTRAL EASTERN DESERT OF EGYPT

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Abstract

The Wadi Hafafit Culmination (WHC) can be subdivided into two main units which are separated by Nugrus thrust shear zone. The WHC domes occupy the southern part of the Central Eastern Desert of Egypt and is cored by five separated gneissic granitoids ranging in composition from tonalite to granodiorite having compositions consistent with hydrous partial melting of a mafic source suggesting subduction-related magmatism. The leucogranites along thrust zones are related to the late phase of metamorphism of Hafafit rocks. The five main domes of the WHC have been labeled A dome occupies the northern part of WHC, B and C domes occupy the central eastern part of WHC, D dome occupies the central western part of WHC and E dome occupies the southern part of WHC. The fourth domes A, B, C and E are affected by more than one deformation event but D dome is affected by one deformation event. The domes are not uniform throughout the WHC according to radial pattern of stretching lineation but these domes of WHC were subjected and affected by late-orogenic extension that was controlled by the Najd transform faults and that resulted in exhumation of the WHC domes are oriented NW-SE parallel to the trend of the left-lateral shear zones of Najd fault system.

Key words: *Wadi Hafafit Culmination, Strain analysis, Deformations, Najd fault.*

1. Introduction

Some of the most conspicuous Precambrian structures in the Eastern Desert of Egypt are met with in the Migif-Hafafit area. This area is characterized by huge open regional folds trending NW-SE. The WHC has been studied by El Ramly et al. (1984) as part of a collaborative project between the Geological Survey of Egypt and University of Mainz, West Germany. This study was based essentially on the detailed lithological field mapping of this area by El Ramly in the late fifties together with photo-interpretation. The most spectacular antiform is a structural culmination composed of several granitoid-cored gneiss domes. These domes are spaced for about 55 km from Wadi Shait in the north to Wadi El Gemal in the south. The WHC is bordered in the east by the Hangaliya-Zabara range and in west by the vast plain of gneissic granitoids which are pierced by conspicuous trachyte plugs (e.g. Gabal Hamr Salma and Gabal El Nuhud). The maximum width of this antiform is 17 km. Hashad (1990) interpreted the granitic gneisses at the core of domes as

gneissic granitic intrusions. They named the rock assemblage above the core granite and below the psammitic gneiss as a metamorphosed and deformed ophiolitic *mélange* assemblage, whereas the psammitic gneiss as a metamorphosed sedimentary unit of a quartzo-feldspathic composition. Several models have been postulated to explain the WHC four tectonic models were proposed to decipher their origin: (1) development of fault-bend fold “antiformal stacks” (e.g. Hafafit domal structure; Greiling et al., 1988a), (2) orogen-parallel crustal extension (e.g. HaFafit, Sibai and Meatiq domal structures; Wallbrecher et al., 1993; Fritz et al., 1996, 2002; Bregar et al., 2002; Loizenbauer et al., 2001; Abdel Wahed, 2008; Khudeir et al., 2008), (3) emplacement within regional domal structures (Ibrahim and Cosgrove, 2001) followed by extension parallel to their fold axes (e.g. Sibai dome, Fowler et al., 2007), and (4) interpreted the WHC as a result of fold interference patterns involving multiply deformed sheath folds (Fowler and El Kalioubi, 2002). The domal structure (WHC) represents the largest antiformal structures in the Nubian Shield. It is considered as one of spectacular structures in the Eastern Desert, but its structural history has not been entirely clarified. The WHC has been subdivided into five separated gneissic domes (labeled A-E core gneisses; of various aerial extends (El Ramly and Greiling, 1988). El Bahariya and Abd El Wahed (2003) considered the granitoid cores together with the genetically related foliated metagabbro to represent the oldest rocks in the Hafafit area. The tectonic evolution of the WHC has been interpreted as a result of fold interference patterns involving multiply deformed sheath folds (Fowler and El Kalioubi, 2002) or as a consequence of tectono-magmatic and tectono-metamorphic processes in an active continental margin tectonic setting (El Bahariya and Abd El Wahed, 2003). The P-T conditions for the different mineral assemblages in order to place some constraints on the petrogenesis of the migmatitic rock association as well as the tectonic processes which accompanied metamorphism and orogenesis (El Bahariya, 2008). The structural characteristics of the northern dome (dome A) of WHC, field descriptions of the dome and surrounding are discussed with emphases to their relations with the overlying Pan-African cover nappes and the bounding ductile shear zones and suggested that the emplacement of gneissic core occurred during accretion of the Pan-African nappes, later strike-slip shear zones of Najd Fault System and the associated Subsidiary shear arrays postdate emplacement of the dome (Shalaby, 2010). Our main aim of this study is to examine the finite strain, variation pattern in gneissic granitoids in five domes of WHC to known tectonic evaluation of this culmination.

2. Geological Setting

The present paper deals with the all granitoids that cored domes of the WHC which is bordered by latitudes 24° 35' 00'' N and 24° 51' 00'' N and longitudes 34° 30' 00'' E and 34° 45' 00'' E. The WHC has been subdivided into five separated gneissic domes (labeled A-E core gneisses; of various aerial extends (El Ramly and Greiling, 1988) (Fig.1a, b). These domes are composed of medium grade gneisses and are separated from the overlying low grade metamorphic rocks by low angle thrust zones. The rock assemblages in Hafafit area could be grouped into two main units which are separated by Nugrus Thrust. The eastern unit (Nugrus unit) is composed mainly of low grade mica-schists and metavolcanics. This unit is associated with remnants of ophiolitic altered ultramafic and metagabbros. The western unit (Hafafit unit) forms Hafafit domes and includes from core to rim. Granite and gneiss of tonalitic and granodiorite composition, banded amphibolites which is overthrust by ultramafic rocks, alternating bands of biotite- and hornblende-gneiss and the psammitic gneiss at the rim of the domal structure. In some parts, the amphibolites are associated with metagabbro (Abd El-Naby and Frisch, 2006). Both units have been intruded by undeformed leucogranites, especially along thrust zones. The present contribution describes the microstructural characteristics of the all from north to south of WHC field descriptions of domes and surrounding with relations of Pan-African and the bounding ductile shear zones. The WHC constitutes an elongated folded belt in Southern Eastern Desert that is about 150 square kilometers, trending NW-SE parallel to the orientation of the Najd fault system. It has evolved and was exhumed in close connection with the activity of this shear zone

system. These granitoids domes are white or greyish white since they consist essentially of oligoclase and quartz with some biotite, hornblende and minor orthoclase. Tonalite prevails in the core of the northernmost dome (dome A), where the peripheral zone is strongly foliated and encloses conformable bands of the foliated metagabbro which several hundreds of meters in length. In places, the tonalite is invaded by numerous, thin pegmatite veinlets giving the erroneous impression of migmatites. The outer rim of this core tonalite is highly injected with white pegmatite veins; it acquires a conspicuous white colour which is accentuated by the weathering and kaolinization of oligoclase. At the northern closure of the dome, highly altered, irregularly-shaped, small masses of rusty brown serpentinite are enclosed in this outer rim, where they are transformed along the peripheries of the intruding pegmatite veins into fibrous anthophyllite (Rasmy, 1974). Stern and Hedge (1985) gave an age of 680 Ma to these tonalites. Whereas the domes B, C and D formed tonalite to granodiorites, the dome E with the conspicuous feature of Naslet Maghar is tonalitic.

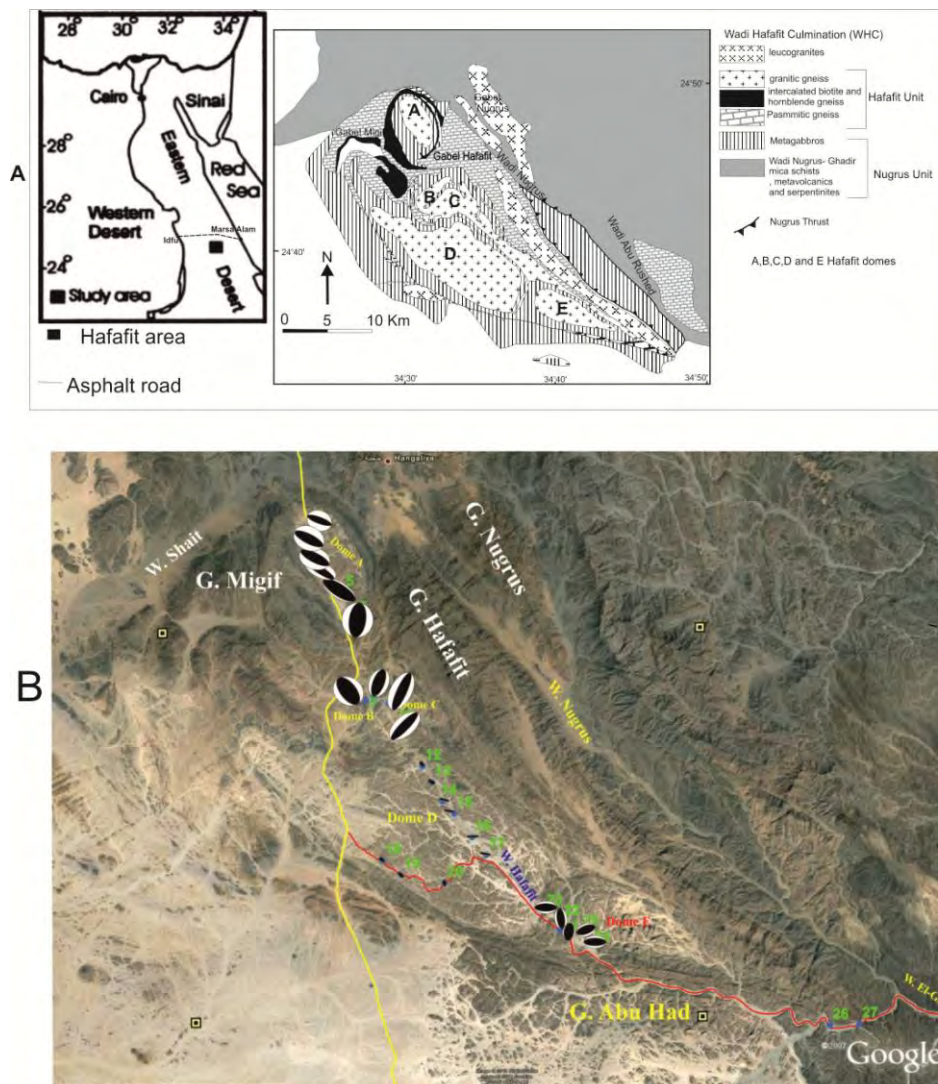


Figure 1 - A- Simplified geological map of Wadi Hafafit area showing major rock types and tectonic structure (modified after El Ramly et al., (1993), B- Photo showing strain ellipsoids in cored granitoids domes of Hafafit.

3. Finite-Strain Analysis

3.1. Field Investigations and Sampling:

24 samples were collected from the Hafafit granitoids domes (for samples localities refer to (Fig. 1b). The Fry and Projection methods used on quartz or feldspar crystal from 6 gneissic tonalite (Dome A), 4 gneissic tonalite and granodiorite (Dome B and C), 9 gneissic tonalite and gneissic granodiorite (Dome D) and 5 gneissic tonalite (Dome E).

Applying the Fry method and projection analysis included: (i) preparation of thin sections along two mutually perpendicular axis (X, Y and Z) subparallel to the XZ and YZ principal planes; (ii) take photographs by polarized microscope; and (iii) employ “fabric 8” program for determining the shape and directions of ellipsoid in each deformation.

3.2 Techniques Used in Strain Analysis:

To quantify the finite strain in the Hafafit granitoids domes, feldspars and quartz have been analysed by Fry and Projection techniques (Fry 1979; Ramsay 1967; Ramsay & Huber 1983; Panozzo, 1984). Two-dimensional strain measurements were made on XY, XZ and YZ sections ($X \geq Y \geq Z$, finite strain axes) in order to estimate the three dimensional strain geometry. The Fry strains are thought to represent the matrix strain, for Fry analysis, the central points of more than 70 feldspar or quartz grains per section were used to calculate strain. The strain estimates were used to calculate the finite-strain ellipsoid according to the modified least-square technique of Owens (1984).

At first the three principal strain directions X, Y and Z have to be determined ($X \geq Y \geq Z$; Principal strain axes). For this purpose, the rock sample is cut parallel to lineation, which represents the plane of flattening (XZ) at the maximum extension direction X, which are normal to the foliation and parallel to lineation (XZ). The intermediate strain direction, Y, is defined by its orthogonality with X and Z, which are normal to foliation and lineation (YZ). After having cut the samples in the Rxz and Ryz directions, Rxy direction has been calculated in accord with the following equation:

$$R_{xy} = R_{xz} / R_{yz}$$

In addition, the stretches S_x , S_y and S_z , which are parallel to principle axes X, Y and Z respectively, enable the following calculation (S) according to (Kassem and Abd El Rahim, 2010).

$$S_y = \sqrt[3]{(R_{xz} / R_{xy}^2)}$$

$$S_x = R_{xy} \times S_y$$

$$S_z = 1 / (S_x \times S_y)$$

3.3 Results of Finite-strain Analysis:

The sample localities for finite strain analysis are shown in Fig. 1 b, the strain data are summarized in tables 1 and 2.

3.3.1 Deformation structures:

In microstructural investigations were made on thin section cut parallel to the foliation (XY), normal to the foliation and parallel to the lineation (XZ) and normal to the foliation and lineation (YZ).

The granitoids domes of the Hafafit is very heterogeneously deformed in the field, where feldspar grains and quartz show different microstructural features in same dome of granite in thin section are extremely smeared out with aspect ratios in XZ and YZ section. The development of normal

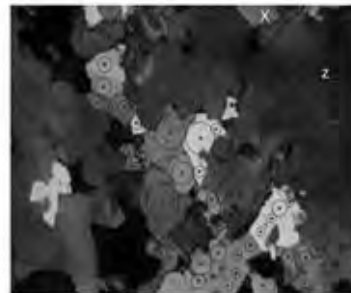
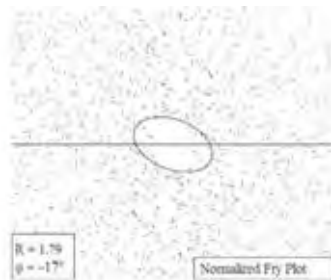
Table 1 – Direction of finite strain axes and stretches for samples from granitoids domes of Gabel Hafafit.

Domes	Rock types	sample no	long.	lat.	X			Y			Z		
					Trend	Plunge	Stretch	Trend	Plunge	Stretch	Trend	Plunge	Stretch
Dome A	Gneissic Tonalite	1BA	34° 30' 09" E	24° 49' 13" N	330	19	1.25	48	43	1.14	310	70	0.70
		2AA	34° 29' 59" E	24° 48' 33" N	330	36	1.35	229	15	1.00	120	50	0.74
		3BA	34° 30' 07" E	24° 48' 18" N	330	31	1.29	7	70	1.07	300	55	0.72
		3CA	34° 30' 07" E	24° 48' 18" N	330	32	1.11	3	69	1.05	302	55	0.86
		4BA	34° 30' 20" E	24° 47' 50" N	330	46	1.25	281	59	1.02	300	40	0.79
Dome B&C	Gneissic Tonalite & Gneissic granodiorite	6AA	34° 37' 05" E	24° 49' 13" N	157	35	1.30	160	73	1.21	125	50	0.63
		7AB	34° 32' 06" E	24° 44' 12" N	150	22	1.20	243	7	1.09	350	67	0.77
		8BB	34° 32' 09" E	24° 44' 06" N	185	25	1.17	92	6	1.12	350	64	0.76
		9AB	34° 32' 06" E	24° 43' 54" N	205	9	1.42	102	51	0.97	290	30	0.72
		10AB	34° 32' 36" E	24° 43' 45" N	225	3	1.20	133	35	1.10	320	55	0.76
Dome D	Gneissic Tonalite & Gneissic granodiorite	12BD	34° 33' 42" E	24° 41' 48" N	150	47	1.15	246	5	1.10	340	43	0.79
		13BD	34° 34' 01" E	24° 41' 18" N	310	41	1.30	340	23	1.08	340	45	0.71
		14BD	34° 34' 22" E	24° 40' 43" N	310	51	1.41	320	15	0.95	340	35	0.75
		15BD	34° 34' 45" E	24° 40' 20" N	280	21	1.41	16	16	1.01	140	63	0.70
		16BD	34° 35' 21" E	24° 39' 36" N	250	0	1.45	160	50	1.12	340	40	0.61
		17BD	34° 35' 49" E	24° 39' 03" N	300	42	1.35	50	21	1.16	160	40	0.64
		18BD	34° 36' 30" E	24° 41' 10" N	310	34	1.12	339	61	1.05	295	55	0.85
		19BD	34° 32' 57" E	24° 38' 28" N	310	70	1.22	49	3	1.03	140	20	0.80
		20BD	34° 34' 24" E	24° 38' 12" N	10	16	1.14	268	36	1.13	120	50	0.78
		21BE	34° 38' 16" E	24° 36' 43" N	325	28	1.44	220	26	1.02	95	50	0.68
Dome E	Gneissic Tonalite	22BE	34° 38' 18" E	24° 36' 58" N	350	10	1.35	250	43	1.08	90	45	0.68
		23BE	34° 37' 43" E	24° 37' 18" N	10	8	1.17	276	29	0.98	115	60	0.87
		24CE	34° 39' 02" E	24° 36' 11" N	250	36	1.23	11	35	1.06	130	35	0.76
		25BE	34° 39' 02" E	24° 36' 24" N	310	38	1.43	212	10	1.07	110	50	0.65

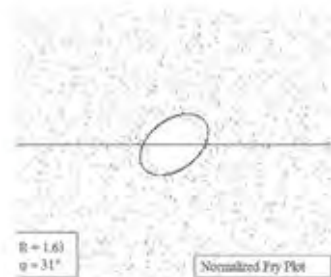
Table 2 – Finite strain data for samples from granitoids domes of Gabel Hafafit.

Hafafit domes	Sample no	Method	R _{xz}	R _{yz}	R _{xy}	Stretch			Principal strain ratio
						S _x (X)	S _y (Y)	S _z (Z)	K
Dome A	1BA	Fry	1.79	1.63	1.098	1.253	1.141	0.700	0.156
	1BA	Projection	1.12	1.07	1.047	1.054	1.007	0.942	0.669
	2AA	Fry	1.83	1.35	1.356	1.354	0.999	0.740	1.016
	2AA	Projection	1.22	1.07	1.140	1.116	0.979	0.915	2.007
	3BA	Fry	1.78	1.48	1.203	1.289	1.072	0.724	0.422
	3BA	Projection	1.18	1.15	1.026	1.066	1.039	0.903	0.174
	3CA	Fry	1.29	1.23	1.049	1.106	1.055	0.857	0.212
	3CA	Projection	1.27	1.05	1.210	1.154	0.954	0.909	4.204
	4BA	Fry	1.59	1.29	1.233	1.251	1.015	0.787	0.802
	4BA	Projection	1.21	1.18	1.025	1.075	1.048	0.888	0.141
	6AA	Fry	2.06	1.91	1.079	1.305	1.210	0.633	0.086
	6AA	Projection	1.22	1.08	1.130	1.113	0.985	0.912	1.623
Dome B&C	7AB	Fry	1.57	1.42	1.106	1.202	1.087	0.765	0.251
	7AB	Projection	1.08	1.08	1.000	1.026	1.026	0.950	0.000
	8BB	Fry	1.54	1.48	1.041	1.170	1.125	0.760	0.084
	8BB	Projection	1.37	1.11	1.234	1.191	0.965	0.870	2.133
	9AB	Fry	1.96	1.34	1.463	1.421	0.971	0.725	1.361
	9AB	Projection	1.32	1.08	1.222	1.173	0.960	0.889	2.778
	10AB	Fry	1.58	1.45	1.090	1.199	1.100	0.759	0.199
	10AB	Projection	1.17	1.15	1.017	1.060	1.042	0.906	0.116
Dome D	12BD	Fry	1.45	1.39	1.043	1.148	1.100	0.792	0.111
	12BD	Projection	1.29	1.14	1.132	1.135	1.003	0.879	0.934
	13BD	Fry	1.82	1.31	1.205	1.299	1.078	0.714	0.403
	13BD	Projection	1.22	1.14	1.070	1.093	1.021	0.896	0.502
	14BD	Fry	1.89	1.27	1.488	1.412	0.949	0.747	1.809
	14BD	Projection	1.15	1.13	1.018	1.054	1.036	0.916	0.136
	15BD	Fry	2	1.43	1.399	1.409	1.007	0.704	0.927
	15BD	Projection	1.31	1.29	1.016	1.100	1.083	0.840	0.053
	16BD	Fry	2.37	1.83	1.295	1.453	1.122	0.614	0.356
	16BD	Projection	1.09	1.04	1.048	1.045	0.997	0.959	1.204
	17BD	Fry	2.1	1.8	1.167	1.348	1.156	0.642	0.208
	17BD	Projection	1.32	1.19	1.109	1.136	1.024	0.860	0.575
	18BD	Fry	1.33	1.24	1.073	1.126	1.050	0.846	0.303
	18BD	Projection	1.14	1.13	1.009	1.048	1.039	0.919	0.068
	19BD	Fry	1.53	1.29	1.186	1.220	1.029	0.796	0.637
	19BD	Projection	1.61	1.4	1.150	1.228	1.068	0.763	0.375
	20BD	Fry	1.46	1.45	1.007	1.137	1.129	0.779	0.015
	20BD	Projection	1.28	1.13	1.133	1.132	0.999	0.884	1.021
Dome E	21BE	Fry	2.12	1.5	1.413	1.442	1.020	0.680	0.827
	21BE	Projection	1.14	1.09	1.046	1.060	1.014	0.930	0.511
	22BE	Fry	1.98	1.59	1.245	1.351	1.085	0.682	0.416
	22BE	Projection	1.12	1.07	1.047	1.054	1.007	0.942	0.669
	23BE	Fry	1.34	1.12	1.196	1.170	0.978	0.873	1.639
	23BE	Projection	1.2	1.08	1.111	1.101	0.991	0.917	1.388
	24CE	Fry	1.61	1.39	1.158	1.231	1.063	0.765	0.406
	24CE	Projection	1.27	1.08	1.176	1.143	0.972	0.900	2.202
	25BE	Fry	2.2	1.64	1.341	1.434	1.069	0.652	0.534
	25BE	Projection	1.24	1.07	1.159	1.128	0.974	0.910	2.275

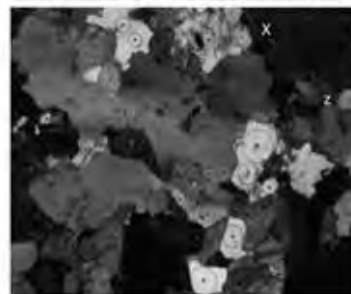
1BAXZ (dome A)



1BAYZ (dome A)



8BBXZ (dome B)



8BBYZ (dome B)



Figure 2 - Fry analysis along XZ and YZ principle finite strain planes and original samples surfaces to which fry applied.

tonalite and granodiorite showing hypidomorphic texture to gneissic tonalite and gneissic granodiorite. In the gneissic granite, the main-phase foliation is a gneissic foliation defined by elongated, dynamically recrystallized feldspar and quartz crystals and alignment of biotite and hornblende (Figs.2, 3).

3.3.2 Magnitudes of Finite Stretches:

The sample localities for finite strain analysis are shown in (Fig. 1 b). The strain data are summarized in tables 1 & 2 shown in a Flinn diagram in (Fig.4), the flinn diagram shows the

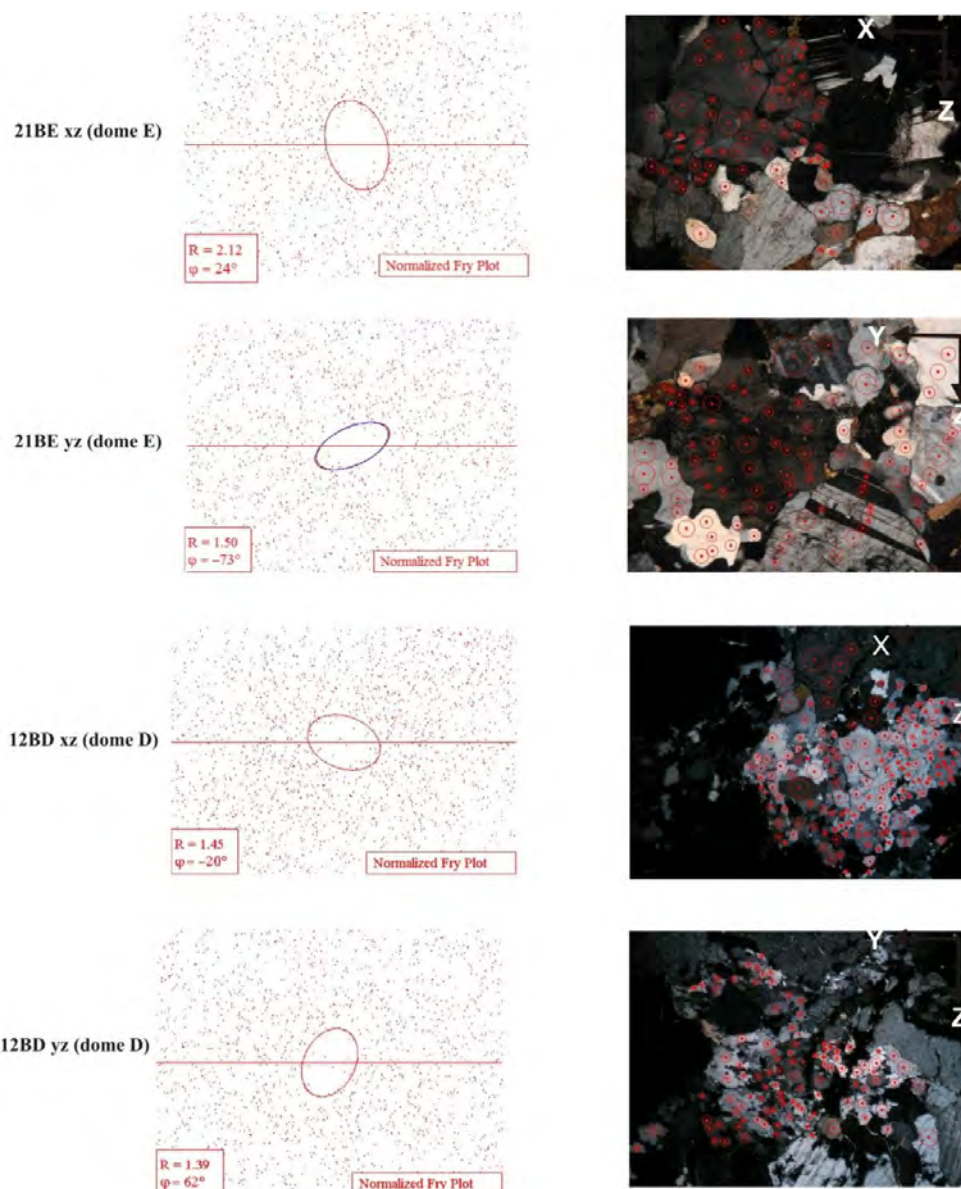


Figure 3 – Fry analysis along XZ and YZ principle finite strain planes and original sample surfaces to which fry applied.

relative shapes of the strain ellipsoids, i.e. prolate vs oblate. To infer strain type, i.e. constrictional vs flattening, information on volume strain is needed.

As shown in Table 2, Projection strains are not fundamentally different from Fry strains; in some cases, the Fry strains are slightly greater than the projection strains. Therefore, our studied samples show that there was no significant difference in deformation behavior between the biotite and Hornblende matrix and the feldspar-quartz porphyroclasts during the accumulation of finite strain during deformation conditions. Furthermore, finite strain in the gneissic tonalite in dome A is of the same order of magnitude as those from the gneissic tonalite to granodiorite of domes B&C and the gneissic tonalite dome E which suggests similar deformation behavior in both lithologies but dome D have another deformation behavior about this domes. In general, the strain ellipsoids in

domes A, B,C and E have oblate and prolate strain symmetry but the strain ellipsoids in dome D have oblate symmetry only. In the dome (A) The axial ratios in XZ sections range from 1.12 to 2.06 with S_x ranging from 1.05 to 1.35. The stretches in the Z direction, S_z range from 0.63 to 0.94 indicating vertical shortening. S_y ranges from 0.95 to 1.20 showing both contraction and extension in this direction.

The strain data verify pronounced heterogeneous deformation of the gneissic tonalite as shown in the field. In dome B&C the axial ratios in XZ sections range from 1.08 to 1.96 with S_x ranging from 1.025 to 1.42. The stretches in the Z direction, S_z range from 0.72 to 0.95 indicating vertical shortening. S_y ranges from 0.95 to 1.12 showing both contraction and extension in this direction. The strain data verify pronounced heterogeneous deformation of the gneissic tonalite to granodiorite as shown in the field. In the dome (E) the axial ratios in XZ sections range from 1.12 to 2.20, with S_x ranging from 1.06 to 1.44. The stretches in the Z direction, S_z range from 0.65 to 0.94 indicating vertical shortening. S_y ranges from 0.97 to 1.08 showing both contraction and extension in this direction. The strain data verify pronounced heterogeneous deformation of the gneissic tonalite as shown in the field. In the dome (D) The axial ratios in XZ sections range from 1.09 to 2.37 with S_x ranging from 1.045 to 1.45. The stretches in the Z direction, S_z range from 0.61 to 0.95 indicating vertical shortening. S_y ranges from 0.99 to 1.20 showing extension in this direction. The strain data verify pronounced homogeneous deformation of the gneissic tonalite and gneissic granodiorite. All domes A, B and C and E of Hafafit granitoids showing flattening and constrictional strains but dome D showing only flattening this means found different events of deformation.

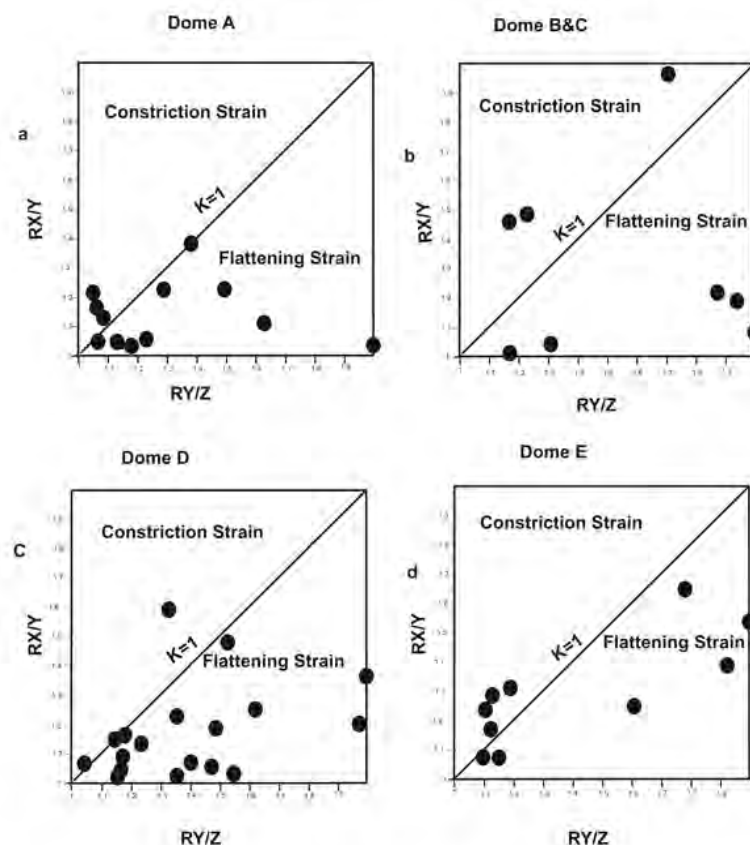


Figure 4 – Flinn diagram showing strain symmetry as obtained by fry and projection methods for Hafafit domes.

5. Discussion

The data show oblate and prolate symmetry (flattening and constriction strains) in gneissic granitoids of domes A, B, C, E but show oblate symmetry (flattening strain) only in granitoids of dome D this indicates that the time of deformation represents the accumulation of ductile to brittle not same time and not same deformation events. This means the granitoids domes of Hafafit are subjected to two different stage of deformation. The early stage of deformation the domes A, B, C, E are affected by more than one deformation event so we observed that the direction of long axis (x) of strain ellipsoid in these gneissic granitoids domes have set W/WSW to N/ENE But granitoids of dome D not affected by this deformation stage. The late stage of deformation all granitoids domes are subjected to general one deformation event so the direction of long axis (x) of strain ellipsoid are common in these domes SE to NW this inductors in cored domes of Hafafit according tectonic evolution.

Structural observations in WHC show that the culmination is bounded from east and west by two non-parallel left-lateral strike-slip shear zones; namely, Nugrus and El Gemal shear zones, where they intersect south of dome "E" by a pivot and spread away northward, giving rise to the V-geometry for the whole culmination (Fig. 5). Published structural data in WHC (e.g. Greiling et al., 1988b; Fritz et al., 1996; Fowler and El Kalioubi, 2002; El Naby et al., 2008; Fowler and Osman, 2009; Shalaby, 2010) elucidate:

1. Northwestward tectonic transportation on thrust zones with gradual decrease of shortening across the culmination from south to north, which is indicated by stronger buckling in dome "E" at the pivot and weak buckling in dome "A" northward. Fabrics in cross-sections parallel to stretching lineation show evidences of northwestward ductile stretching involved within gneissic bands especially in domes located further north.
2. Opposite to pivot and especially north of zone between domes "A" and "B", ductile normal shears are frequent, dipping mostly shallowly to NW. These extensional fabrics are reactivated into brittle normal faults of regional extend. (e.g. Shait normal fault)
3. The northern area is intensively crossed by roughly E–W trending dykes that rather disappear southward.
4. The dome "B" and "C" is crossed by N–S trending left-lateral strike-slip fault. Displacement on this fault is fading towards zero at dome "D".
5. Late –orogenic extension and crustal thinning that controlled by the Najd transform faults that resulted in exhumation of the Hafafit domes through a combination of transpression and lateral extrusion.

6. Conclusions

The strain analysis of the deformation in WHC domes south central Eastern Desert of Egypt indicates that ductile deformation show oblate and prolate symmetry (flattening and constriction strains) in gneissic granitoids of WHC domes A, B, C, E and show oblate symmetry (flattening strain) only in granitoids of dome D. The oblate symmetry (flattening) on all domes due to late orogenic extension and crustal thinning associated the Najd transform fault that resulted exhumation of the WHC domes, The domes A, B, C and E show prolate symmetry (constriction strains for this reasons:

1. Dome A sheared with Shait-Nugrus Shear zone (SNSZ) after post arc collision structure (Fowler and Osman, 2009).
2. Dome B & C shows constriction strain because more complex patterns of superimposed interference folds and sheared with Shait-Nugrus Shear zone (SNSZ) (Fowler and El Kalioubi, 2002).
3. Dome E is indicated by stronger buckling in dome "E" at the pivot of two shear zones (Nugrus and El Gemal shear zones).

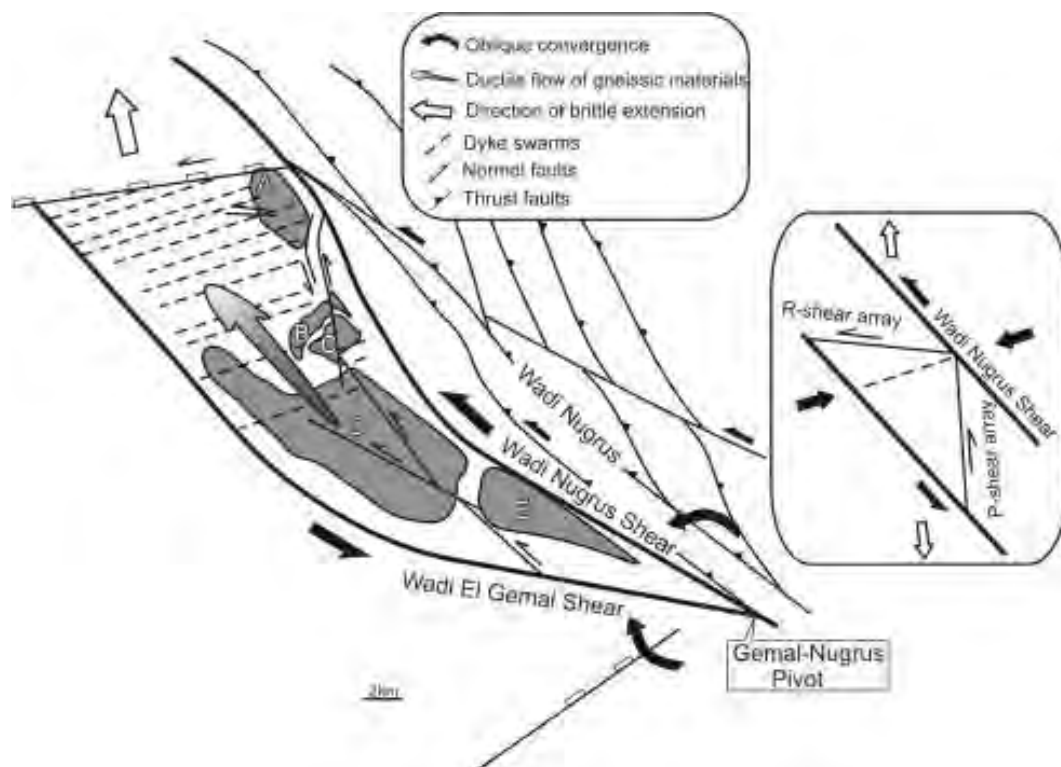


Figure 5 - Simplified structural map for the WHC, collected from landsat images and different literatures (e.g. Greiling et al., 1988b and Fritz et al., 1996; Fowler and El Kalioubi, 2002, Fowler and Osman, 2009 and Shalaby, 2010) The V-geometry for the whole culmination is formed with intersection of two non-parallel left-lateral strike-slip shear zones (Nugrus and El Gemal shear zones) at a pivot, located south of dome “E”, while they spread away northward. Stronger domal buckling close to the pivot and weak buckling further north, north-westward ductile stretching in gneissic domes located further north with development of early ductile, later brittle, Sha’it normal faults, the frequent distribution of E–W trending dykes to north and their rather disappear southward, are structural elements that satisfy the synthetic component of the scissor-like deformation model described by Fowler and Osman (2001). In this model, exhumation of Hafafit gneisses is expected to be accomplished by Late –orogenic extension and crustal thinning that controlled by the Najd transform faults that resulted in exhumation of the Hafafit domes through a combination of transpression and lateral extrusion.

7. Acknowledgments

Our thanks are due to Greece Scholarships Foundation (IKY) for supporting the doctoral studies in Aristotle University Thessaloniki (Greece) our thanks are also Professor Mohamed Abdel Wahed (Tanta University) for helping me in field work.

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